

Numerical simulation of soil water dynamics: transient flow

The relation between ψ_m and the unsaturated hydraulic conductivity $k(\psi_m)$ or the differential soil water capacity $C(\psi_m)$ causes the differential equation (DE) to be strongly non-linear!

- ➔ There are no analytical solutions available to solve the DE for initial and boundary conditions found under **field conditions**.
- ➔ So: numerical solutions will be needed to solve the DE:
 - finite difference solutions of the DE
 - finite element solutions of the DE

There are different methods to solve the non-linear DE, example 1D-vertical:

1. Use time steps Δt and depth compartments Δz small enough so $k(\psi_m)$ and $C(\psi_m)$ can be assumed constant during the time step Δt (explicit and implicit numerical solutions)

Finite-difference methods

partial differential equations

$$\frac{\delta \theta}{\delta t} = \frac{-\delta q}{\delta z} = \frac{\delta}{\delta z} \left(k(\theta) \cdot \left(\frac{\delta \psi}{\delta z} + 1 \right) \right) - S$$

$$\frac{\delta \psi}{\delta t} = \frac{1}{C(\psi)} \cdot \frac{\delta}{\delta z} \left(k(\psi) \cdot \left(\frac{\delta \psi}{\delta z} + 1 \right) \right) - S$$

→ Transform into finite-difference equations

Time and space discretisation: time steps Δt and depth compartments Δz

Approximate terms in



Taylor theorem:

$$\psi(t + \Delta t) = \psi(t) + \Delta t \cdot \psi'(t) + \frac{1}{2} \Delta t^2 \cdot \psi''(t) + \dots$$

$$\theta(t + \Delta t) = \theta(t) + \Delta t \cdot \theta'(t) + \frac{1}{2} \Delta t^2 \cdot \theta''(t) + \dots$$

Forward finite-difference quotient ψ

$$\frac{\delta\psi}{\delta t} \approx \frac{\psi(t + \Delta t) - \psi(t)}{\Delta t} \quad O(\Delta t)$$

Forward finite-difference quotient θ

$$\frac{\delta\theta}{\delta t} \approx \frac{\theta(t + \Delta t) - \theta(t)}{\Delta t} \quad O(\Delta t)$$

- assuming second and higher powers of Δt can be neglected \rightarrow errors are of order $O(\Delta t)$
- thus : numerical solutions are an approximation of the true solution due to the neglect of rest-terms

Central differential quotient:

$$\psi(z + \Delta z) = \psi(z) + \Delta z \cdot \psi'(z) + \frac{1}{2} \Delta z^2 \cdot \psi''(z) + \dots \quad (1)$$

$$\psi(z - \Delta z) = \psi(z) - \Delta z \cdot \psi'(z) + \frac{1}{2} \Delta z^2 \cdot \psi''(z) - \dots \quad (2)$$

(1) - (2):

$$\psi'(z) = \frac{\delta\psi}{\delta z} \approx \frac{\psi(z + \Delta z) - \psi(z - \Delta z)}{2\Delta z} \quad \begin{array}{l} O(\Delta z^2) \\ (= \text{error}) \end{array}$$

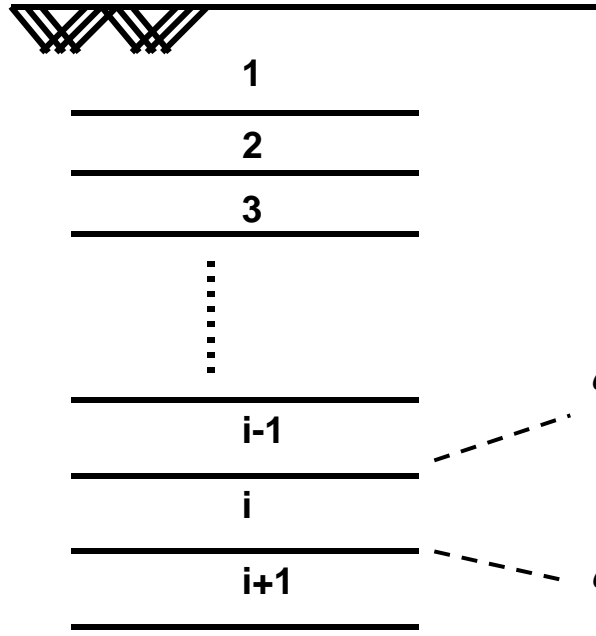
(1) + (2):

$$\psi''(z) = \frac{\delta^2\psi}{\delta z^2} \approx \frac{\psi(z + \Delta z) - 2\psi(z) + \psi(z - \Delta z)}{\Delta z^2} \quad \begin{array}{l} O(\Delta z^3) \\ (= \text{error}) \end{array}$$

Error: - order O of the approximation
- approximation in time and space due to discretisation

Explicit scheme with explicit linearization:

Spatial discretization:



$$q_{i-\frac{1}{2}}^j = -k_{i-\frac{1}{2}}^j \left(\frac{\psi_{i-1}^j - \psi_i^j}{\Delta z} + 1 \right)$$

$$\psi = \psi_m$$

$$q_{i+\frac{1}{2}}^j = -k_{i+\frac{1}{2}}^j \left(\frac{\psi_i^j - \psi_{i+1}^j}{\Delta z} + 1 \right)$$

$$\frac{\delta \theta}{\delta t} = -\frac{\delta q}{\delta z} = \frac{\delta}{\delta z} \left[k(\psi) \cdot \left(\frac{\delta \psi_m}{\delta z} + 1 \right) \right]$$

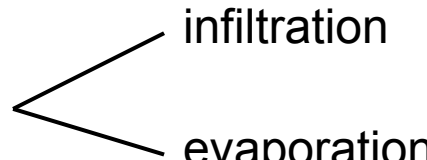
$$\frac{\theta_i^{j+1} - \theta_i^j}{\Delta t} = \frac{1}{\Delta z} \left[k_{i-\frac{1}{2}}^j \cdot \frac{(\psi_{i-1}^j - \psi_i^j)}{\Delta z} + k_{i-\frac{1}{2}}^j \right] - \frac{1}{\Delta z} \left[k_{i+\frac{1}{2}}^j \cdot \frac{(\psi_i^j - \psi_{i+1}^j)}{\Delta z} + k_{i+\frac{1}{2}}^j \right]$$

rearrange:

$$\theta_i^{j+1} = \theta_i^j + \frac{\Delta t}{\Delta z} \left[k_{i-\frac{1}{2}}^j \cdot \frac{(\psi_{i-1}^j - \psi_i^j)}{\Delta z} + k_{i-\frac{1}{2}}^j \right] - \frac{\Delta t}{\Delta z} \left[k_{i+\frac{1}{2}}^j \cdot \frac{(\psi_i^j - \psi_{i+1}^j)}{\Delta z} + k_{i+\frac{1}{2}}^j \right]$$

$$\left. \begin{aligned} k_{i-\frac{1}{2}}^j &= \sqrt{k(\psi_{i-1}^j) \cdot k(\psi_i^j)} \\ k_{i+\frac{1}{2}}^j &= \sqrt{k(\psi_i^j) \cdot k(\psi_{i+1}^j)} \end{aligned} \right\} \text{geometrical mean}$$

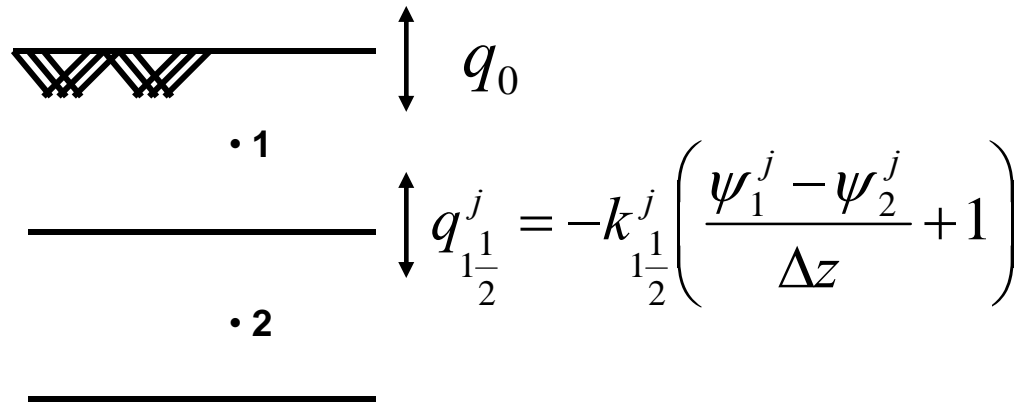
Boundary conditions:

- at the top \rightarrow flux q_0 

- at the bottom \rightarrow groundwater table $\psi_m = 0$

\rightarrow prescribe ψ_m e.g. $\psi_m = -50\text{cm}$

The solution of the top nodal point:

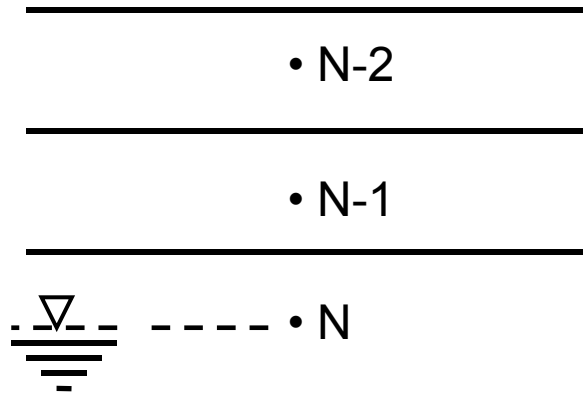


$q_0 < 0$ infiltration

$q_0 > 0$ evaporation

$$\theta_1^{j+1} = \theta_1^j - \frac{\Delta t}{\Delta z} q_0 - \frac{\Delta t}{\Delta z} \left[k_{1/2}^j \left(\frac{\psi_1^j - \psi_2^j}{\Delta z} + 1 \right) \right]$$

Solution for the bottom boundary condition



ψ_N is known, so N-1 is the last node for which the equation must be solved.

$$\theta_{N-1}^{j+1} = \theta_{N-1}^j + \frac{\Delta t}{\Delta z} \left[k_{N-1\frac{1}{2}}^j \left(\frac{\psi_{N-2}^j - \psi_{N-1}^j}{\Delta z} + 1 \right) \right] - \frac{\Delta t}{\Delta z} \left[k_{N-\frac{1}{2}}^j \left(\frac{\psi_{N-1}^j - \psi_N}{\Delta z} + 1 \right) \right]$$

Stability and convergence

Explicit linearization means that Δz and Δt should be taken small enough to secure an accurate numerical solution.

- time step Δt : Δt is chosen so, that the water content change $\Delta\theta$ of any compartment in 1 timestep is less than $0.001 \text{ (cm}^3/\text{cm}^3\text{)}$

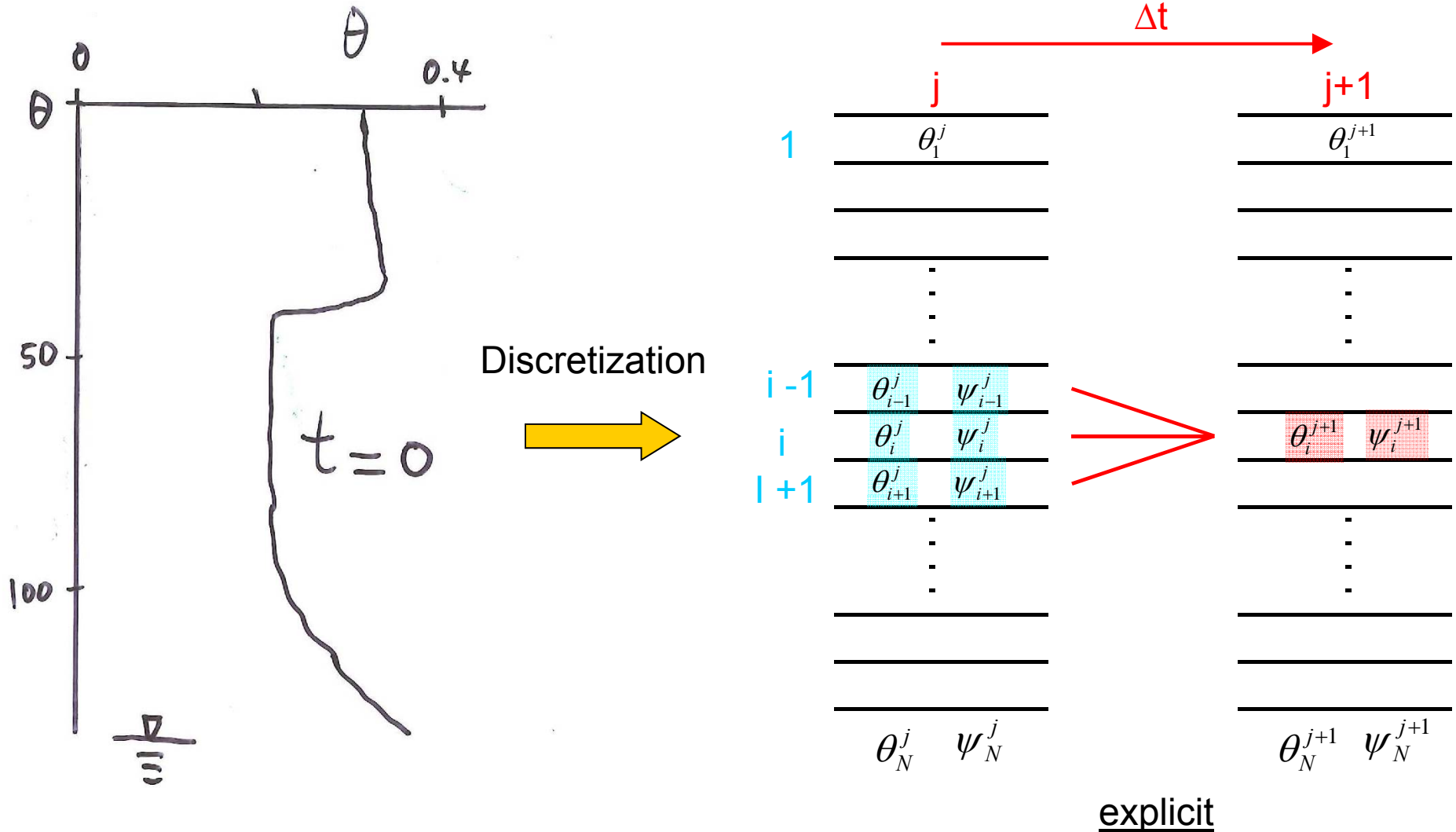
thus:
$$\Delta t = \frac{\Delta\theta \cdot \Delta z}{\left| \Delta q_i^n \right|}$$
 with $\left| \Delta q_i^n \right| = \text{netto flux}$

- space increment Δz : depends on how accurate the solution is wanted. A smaller Δz gives a better resolution and solution

normally: $1 < \Delta z < 10 \text{ cm}$

in parts of the profile where the changes of θ and ψ are small, Δz can be larger (deeper parts of the profile)

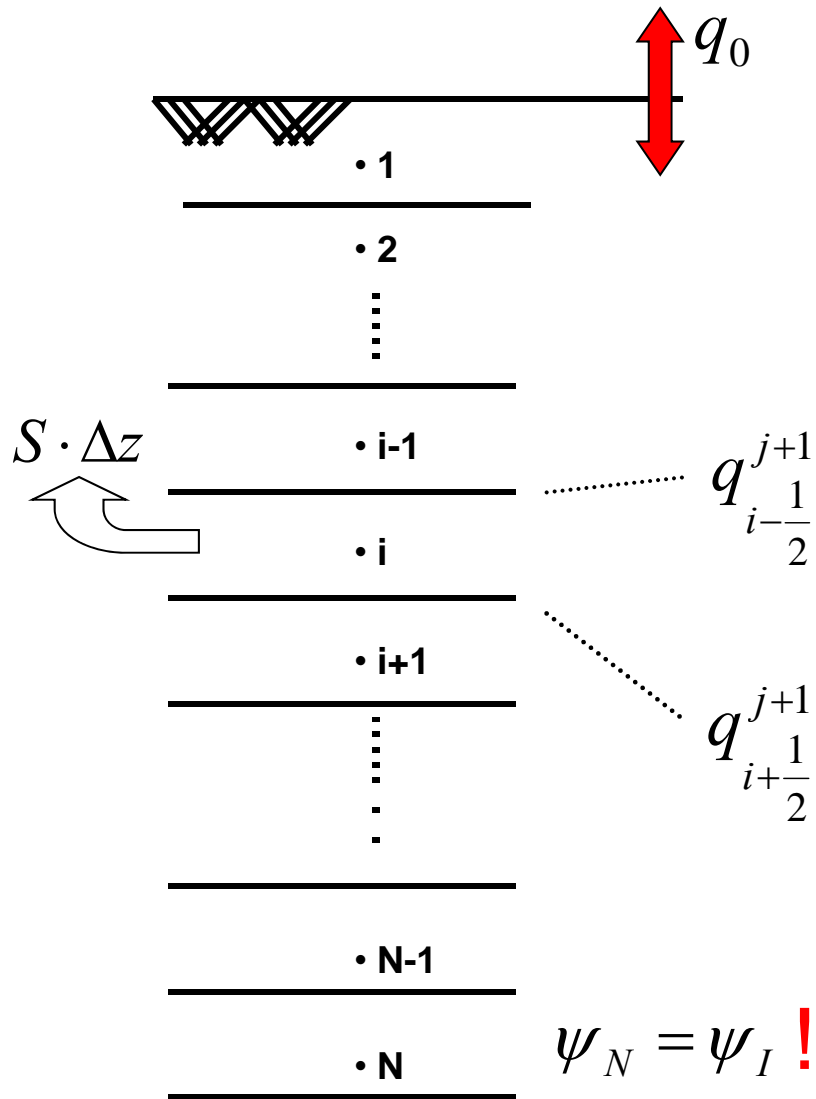
Numerical Model: explicit scheme



Boundary conditions:

- $z = 0$ and $t \geq 0$: rainfall and evaporation
- $z > 0$ and $t \geq 0$: root water uptake $S(z,t)$
- $z_N(t)$ and $t \geq 0$: e.g. given water table depth $z_N(t)$

Spatial discretization:



Implicit scheme with explicit linearization:

$$\frac{\delta \psi}{\delta t} = \frac{I}{C(\psi)} \cdot \frac{\delta}{\delta z} \left[k(\psi) \left(\frac{\delta \psi}{\delta z} + 1 \right) \right] - \frac{S}{C(\psi)}$$

$$\psi = \psi_m$$

$$q_{i-\frac{1}{2}}^{j+1} = -k_{i-\frac{1}{2}}^j \left(\frac{\psi_{i-1}^{j+1} - \psi_i^{j+1}}{\Delta z} + 1 \right)$$

$$q_{i+\frac{1}{2}}^{j+1} = -k_{i+\frac{1}{2}}^j \left(\frac{\psi_i^{j+1} - \psi_{i+1}^{j+1}}{\Delta z} + 1 \right)$$

$$\psi_N = \psi_I !$$

$$\frac{\psi_i^{j+1} - \psi_i^j}{\Delta t} = \frac{1}{C_i^j \Delta z} \left[k_{i-\frac{1}{2}}^j \left(\frac{\psi_{i-1}^{j+1} - \psi_i^{j+1}}{\Delta z} + 1 \right) - k_{i+\frac{1}{2}}^j \left(\frac{\psi_i^{j+1} - \psi_{i+1}^{j+1}}{\Delta z} + 1 \right) \right] - \frac{S_i}{C_i^j}$$

rearrange:

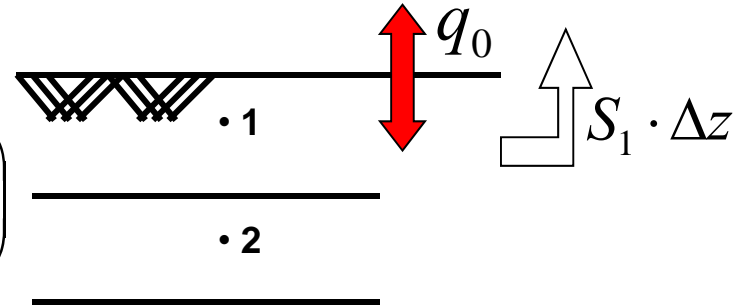
$$-\frac{\Delta t \cdot k_{i-\frac{1}{2}}^j}{C_i^j \Delta z^2} \psi_{i-1}^{j+1} + \left(1 + \frac{\Delta t \cdot k_{i-\frac{1}{2}}^j}{C_i^j \Delta z^2} + \frac{\Delta t \cdot k_{i+\frac{1}{2}}^j}{C_i^j \Delta z^2} \right) \psi_i^{j+1} - \frac{\Delta t \cdot k_{i+\frac{1}{2}}^j}{C_i^j \Delta z^2} \psi_{i+1}^{j+1} =$$

$$\psi_i^j + \frac{\Delta t}{C_i^j \Delta z} \left(k_{i-\frac{1}{2}}^j - k_{i+\frac{1}{2}}^j \right) - \frac{S_i \Delta t}{C_i^j}$$

$$\Rightarrow -A_i \psi_{i-1}^{j+1} + B_i \psi_i^{j+1} - C_i \psi_{i+1}^{j+1} = D_i$$

Upper boundary condition: solution for the 1th node

$$q_{1\frac{1}{2}}^{j+1} = -k_{1\frac{1}{2}}^j \left(\frac{\psi_1^{j+1} - \psi_2^{j+1}}{\Delta z} + 1 \right)$$



$$\frac{\psi_1^{j+1} - \psi_1^j}{\Delta t} = \frac{q_0}{C_1^j \Delta z} - \frac{\Delta t}{C_1^j \Delta z} \left[k_{1\frac{1}{2}}^j \left(\frac{\psi_1^{j+1} - \psi_2^{j+1}}{\Delta z} + 1 \right) \right] - \frac{S_1}{C_1^j}$$

rearrange:

$$+ \left(1 + \frac{\Delta t \cdot k_{1\frac{1}{2}}^j}{C_1^j \Delta z^2} \right) \psi_1^{j+1} - \left(\frac{\Delta t \cdot k_{1\frac{1}{2}}^j}{C_1^j \Delta z^2} \right) \psi_2^{j+1} = \psi_1^j + \frac{\Delta t \cdot q_0}{C_1^j \Delta z} - \frac{\Delta t \cdot k_{1\frac{1}{2}}^j}{C_1^j \Delta z} - \frac{S_1 \Delta t}{C_1^j}$$

$$\Rightarrow + B_1 \psi_1^{j+1} - C_1 \psi_2^{j+1} = D_1$$

Lower boundary condition: solution for the node N-1

$\Rightarrow \psi_{N-i}^{j+i}$ is the last unknown !

• N-2

• N-1

• N $\psi_N^j = \psi_N^{j+1} = \psi_N$

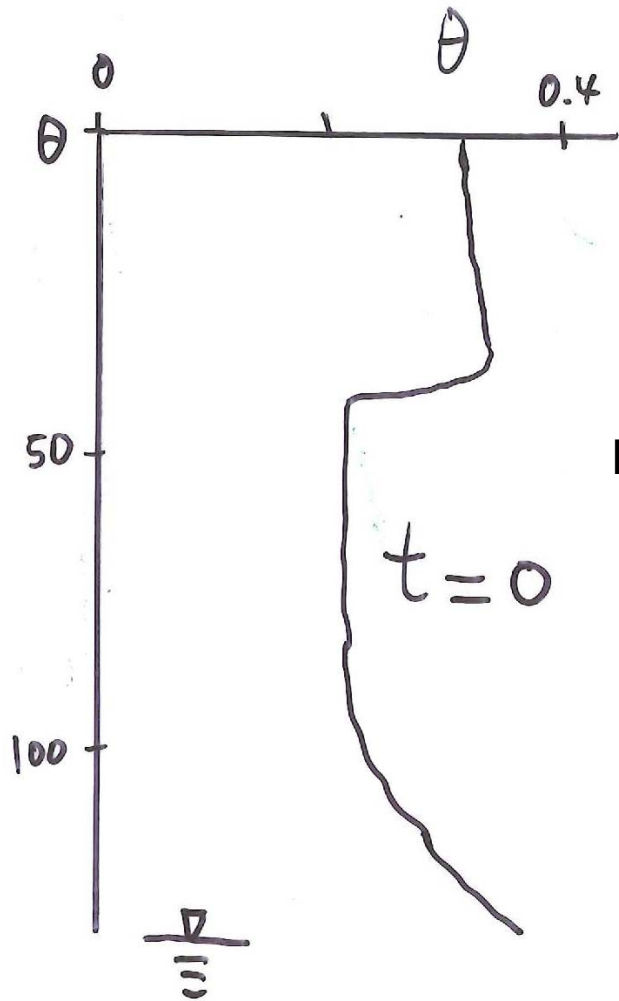
$$\frac{\psi_{N-1}^{j+1} - \psi_{N-1}^j}{\Delta t} = \frac{1}{C_{N-1}^j \Delta z} \left[k_{N-1\frac{1}{2}}^j \left(\frac{\psi_{N-2}^{j+1} - \psi_{N-1}^{j+1}}{\Delta z} + 1 \right) - k_{N-1\frac{1}{2}}^j \left(\frac{\psi_{N-1}^{j+1} - \psi_N^{j+1}}{\Delta z} + 1 \right) \right] - \frac{S_{N-1}}{C_{N-1}^j}$$

rearrange:

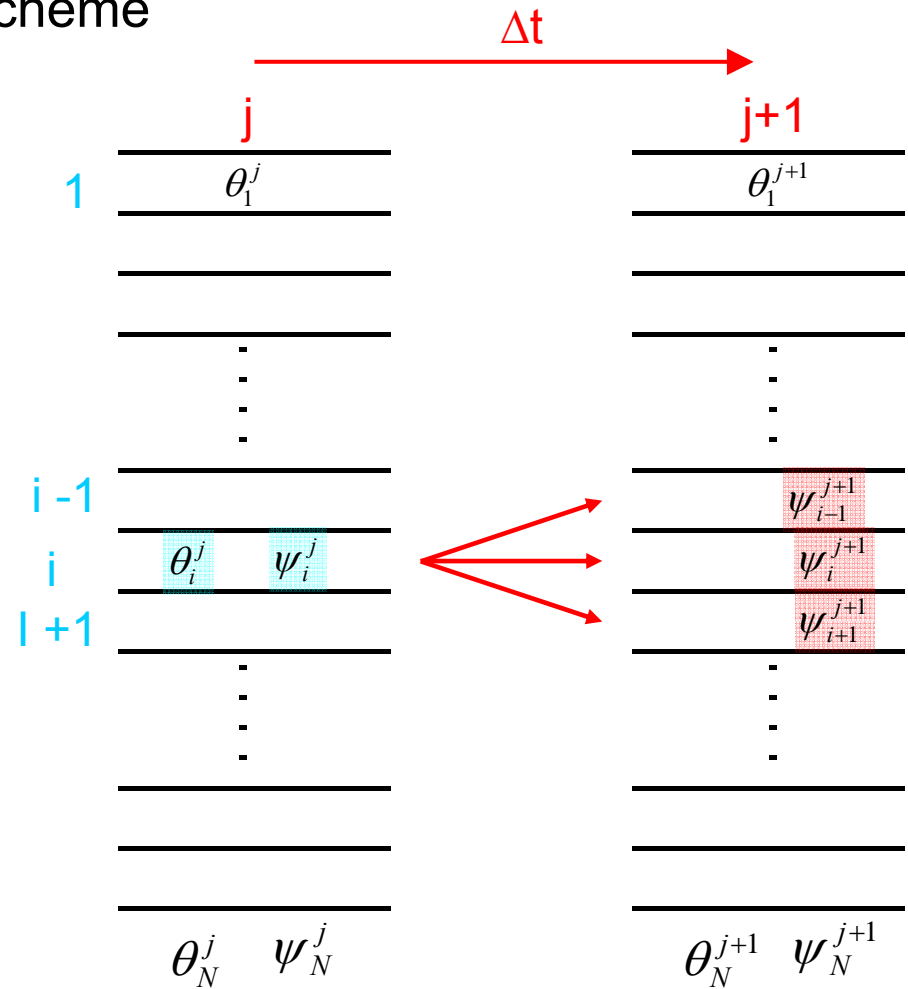
$$-\left(\frac{\Delta t \cdot k_{N-1\frac{1}{2}}^j}{C_{N-1}^j \Delta z^2} \right) \psi_{N-2}^{j+1} + \left(1 + \frac{\Delta t \cdot k_{N-1\frac{1}{2}}^j}{C_{N-1}^j \Delta z^2} + \frac{\Delta t \cdot k_{N-1\frac{1}{2}}^j}{C_{N-1}^j \Delta z^2} \right) \psi_{N-1}^{j+1} = \psi_{N-1}^j + \left(\frac{\Delta t \cdot k_{N-1\frac{1}{2}}^j}{C_{N-1}^j \Delta z^2} \right) \psi_N^{j+1} + \frac{\Delta t}{C_{N-1}^j \Delta z} \left(k_{N-1\frac{1}{2}}^j - k_{N-1\frac{1}{2}}^j \right) - \frac{S_{N-1} \Delta t}{C_{N-1}^j}$$

$$\Rightarrow -A_{N-1} \psi_{N-2}^{j+1} + B_{N-1} \psi_{N-1}^{j+1} = D_{N-1}$$

Numerical Model: Implicit scheme



Discretization



Initial conditions:

- $\theta(z, t=0)$
- $\psi(z, t=0)$

Boundary conditions:

- $z = 0$ and $t \geq 0$: Rainfall N and Evaporation E_a
- $z > 0$ and $t \geq 0$: root water uptake $S(z, t)$
- $z_N(t)$ and $t \geq 0$: e.g. given water table depth $z_N(t)$

Solving systems of linear equations:

$$\frac{\psi_i^{j+1} - \psi_i^j}{\Delta t} = \frac{1}{C_i^j \Delta t} \left\{ k^{j, i-\frac{1}{2}} \left(\frac{\psi_{i-1}^{j+1} - \psi_i^{j+1}}{\Delta z} + 1 \right) - k^{j, i+\frac{1}{2}} \left(\frac{\psi_i^{j+1} - \psi_{i+1}^{j+1}}{\Delta z} + 1 \right) \right\}$$

1. Iterative Methods:

e.g. Gauss-Seidel
iteration scheme

$$-A_i \psi_{i-1}^{j+1} + B_i \psi_i^{j+1} - C_i \psi_{i+1}^{j+1} = D_i$$

General iteration eq. :

$$[\psi_i^{j+1}]^{m+1} = \frac{A_i}{B_i} [\psi_{i-1}^{j+1}]^{m+1} + \frac{C_i}{B_i} [\psi_{i+1}^{j+1}]^m + \frac{D_i}{B_i}$$

$m, m+1 = m^{\text{th}}$ and $(m+1)^{\text{th}}$ approximation of ψ_i^{j+1}

if $\left| [\psi_i^{j+1}]^{m+1} - [\psi_i^{j+1}]^m \right| < \varepsilon$ is satisfied for all $i = 1, 2, \dots, n$



sufficient approximation

1. node:

$$+ B_1 \psi_1^{j+1} - C_1 \psi_2^{j+1} = D_1$$

$$(\psi_1^{j+1})^{m+1} = \left(\frac{C_1}{B_1} (\psi_2^{j+1})^m + \frac{D_1}{B_1} \right)$$

1. Iteration m=1:

$$[\psi_1^{j+1}]^2 = \left(\frac{C_1}{B_1} [\psi_2^{j+1}]^1 + \frac{D_1}{B_1} \right)$$

2 – (N-1) nodes

$$- A_i \psi_{i-1}^{j+1} + B_i \psi_i^{j+1} - C_i \psi_{i+1}^{j+1} = D_i$$

$$[\psi_i^{j+1}]^{m+1} = \frac{A_i}{B_i} [\psi_{i-1}^{j+1}]^{m+1} + \frac{C_i}{B_i} [\psi_{i+1}^{j+1}]^m + \frac{D_i}{B_i}$$

1. Iteration m=1:

$$[\psi_i^{j+1}]^2 = \frac{A_i}{B_i} [\psi_{i-1}^{j+1}]^2 + \frac{C_i}{B_i} [\psi_{i+1}^{j+1}]^1 + \frac{D_i}{B_i}$$

2 – (N-1) nodes

$$-A_i \psi_{i-1}^{j+1} + B_i \psi_i^{j+1} - C_i \psi_{i+1}^{j+1} = D_i$$

$$\left[\psi_i^{j+1} \right]^{m+1} = \frac{A_i}{B_i} \left[\psi_{i-1}^{j+1} \right]^{m+1} + \frac{C_i}{B_i} \left[\psi_{i+1}^{j+1} \right]^m + \frac{D_i}{B_i}$$

1. Iteration $m=4$:

$$\left[\psi_i^{j+1} \right]^5 = \frac{A_i}{B_i} \left[\psi_{i-1}^{j+1} \right]^5 + \frac{C_i}{B_i} \left[\psi_{i+1}^{j+1} \right]^4 + \frac{D_i}{B_i}$$

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$$\left| \left[\psi_i^{j+1} \right]^{m+1} - \left[\psi_i^{j+1} \right]^m \right| < \varepsilon \quad \text{is satisfied for all } i = 1, 2, \dots, n$$



sufficient approximation

2. Direct methods: e.g. Gauss elimination

If we have $N-1$ internal mesh points, we have the following system of linear equations:

$$+ b_1 u_1 - c_1 u_2 = d_1 \quad (1) \quad u_1 = \psi_1^{j+1}$$

$$- a_2 u_1 + b_2 u_2 - c_2 u_3 = d_2 \quad (2)$$

$$- a_i u_{i-2} + b_i u_i - c_i u_{i+1} = d_i \quad (i)$$

$$- a_{N-1} u_{N-2} + b_{N-1} u_{N-1} = d_{N-1} \quad (N-1)$$

- a's, b's, c's and d's are known
- elimination: u1 from eq.2
u2 from eq.3
etc.

Assume the following stage of elimination:

$$a_{i-1}u_{i-1} - c_{i-1}u_i = S_{i-1}$$

$$-a_i u_{i-1} + b_i u_i - c_i u_{i+1} = d_i$$

Eliminating u_{i-1} leads to:

$$\left(b_i - \frac{a_i c_{i-1}}{a_{i-1}} \right) u_i - c_i u_{i+1} = d_i + \frac{a_i S_{i-1}}{a_{i-1}}$$

$$\text{or } \alpha_i u_i - c_i u_{i+1} = S_i$$

$$\text{with } \alpha_i = b_i - \frac{a_i c_{i-1}}{a_{i-1}}$$

$$\alpha_1 = b_1$$

$$\text{and } S_i = d_i + \frac{a_i S_{i-1}}{a_{i-1}}$$

$$S_1 = d_1$$

The last pair of simultaneous equations are:

$$\alpha_{N-2}u_{N-2} - c_{N-2}u_{N-1} = S_{N-2}$$

and

$$-a_{N-1}u_{N-2} - b_{N-1}u_{N-1} = d_{N-1}$$

Elimination of u_{N-2} leads to:

$$\left[b_{N-1} - \frac{a_{N-1}c_{N-2}}{\alpha_{N-2}} \right] u_{N-1} = d_{N-1} + \frac{a_{N-1}S_{N-2}}{\alpha_{N-2}}$$

or:

$$\alpha_{N-1}u_{N-1} = S_{N-1}$$

By backward substitution we will find the unknown u_i :



$$u_{N-1} = \frac{S_{N-1}}{\alpha_{N-1}}$$

$$u_i = \frac{1}{\alpha_i} (S_i + c_i u_{i+1}) \quad i=N-2, N-3, \dots, 1$$

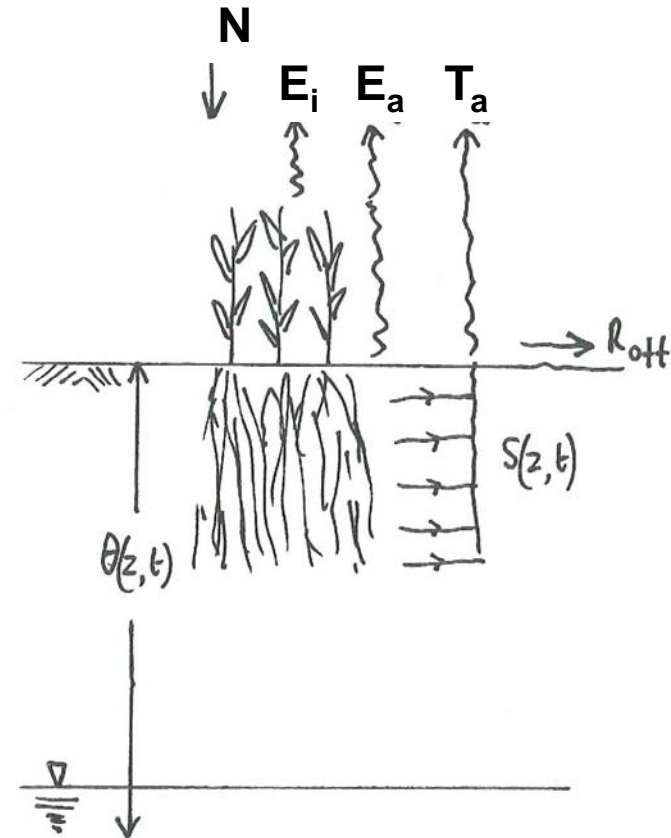
Time-depth-curves:

a method to visualize the vertical movement of water in the soil

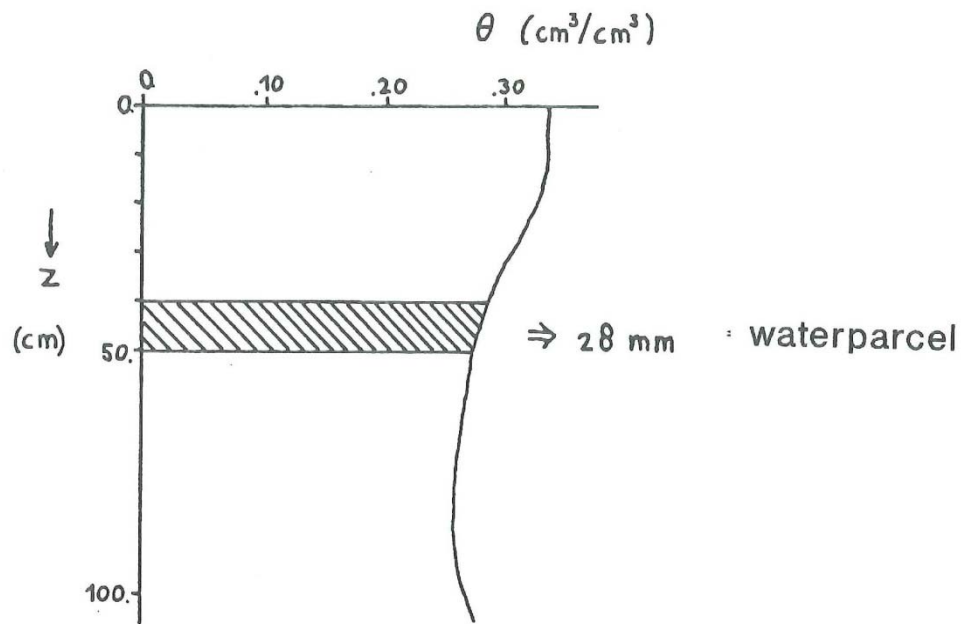
Data needed to : - generate time-depth-curves
- calculate the water balance
of water-parcels between
time-depth-curves

are:

- Precipitation N (cm/d)
- Interception E_i (cm/d)
- actual Evaporation E_a (cm/d)
- root water uptake as a function of
time and depth $S(z,t)$ ($\text{cm}^3/\text{cm}^3 \cdot \text{d}$)
- volumetric water content as a function
of time and depth $\theta(z,t)$ ($\text{cm}^3/\text{cm}^3 \cdot \text{d}$)
- surface runoff R_{off} (cm/d)



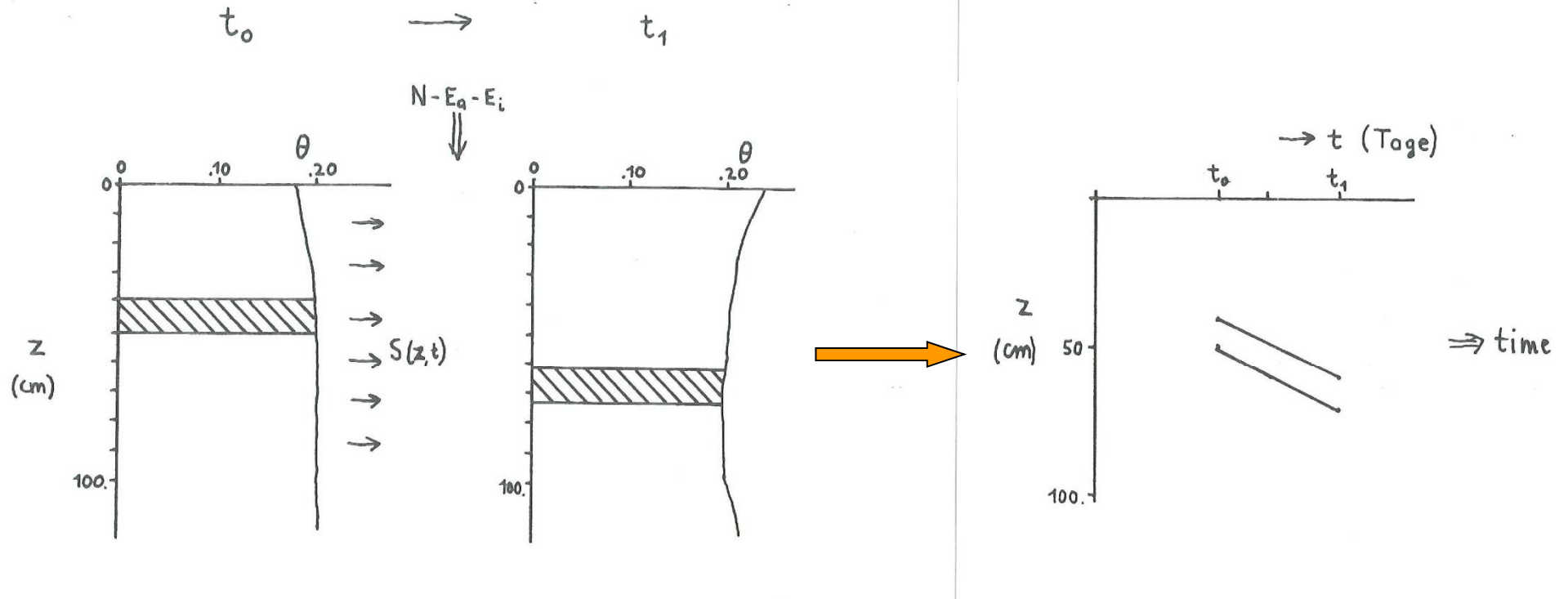
initialize waterparcels at t=0



assumptions: only piston flow
no diffusion-dispersion

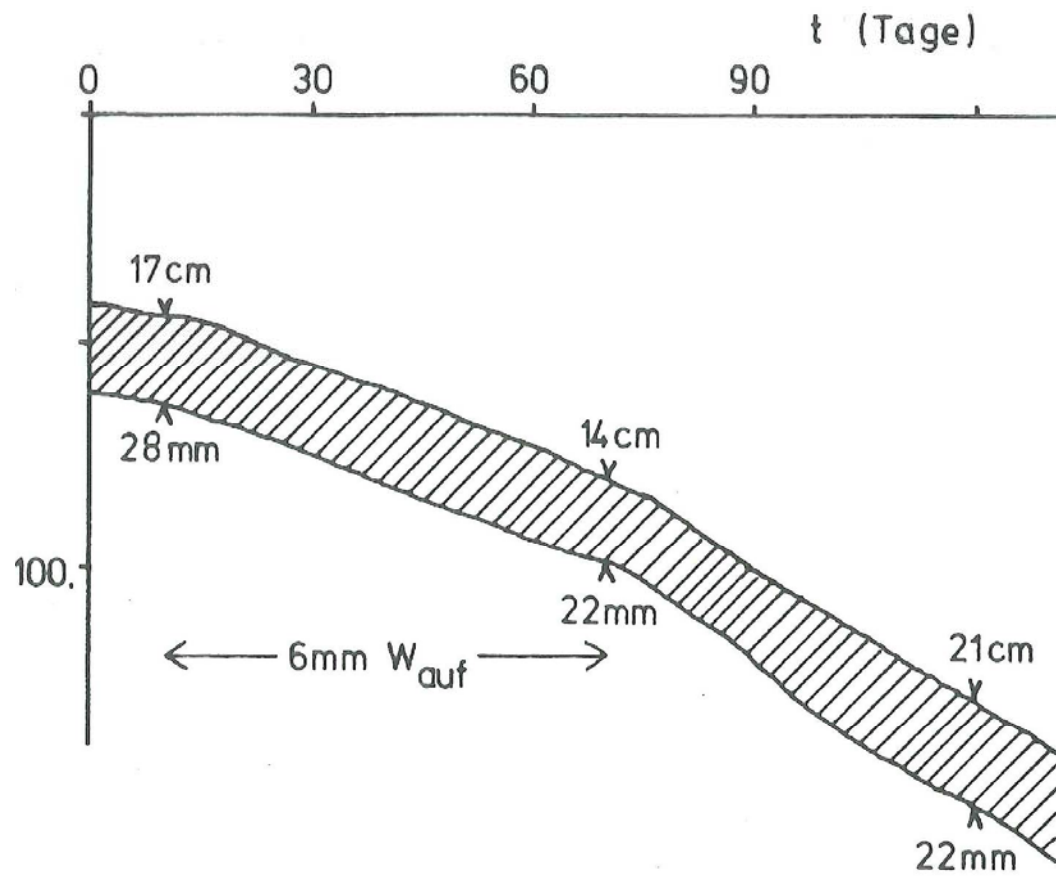
$$W = \int_{z1=40}^{z2=68} \theta(z,0) dz = 28 \text{ mm}$$

movement of a waterparcel

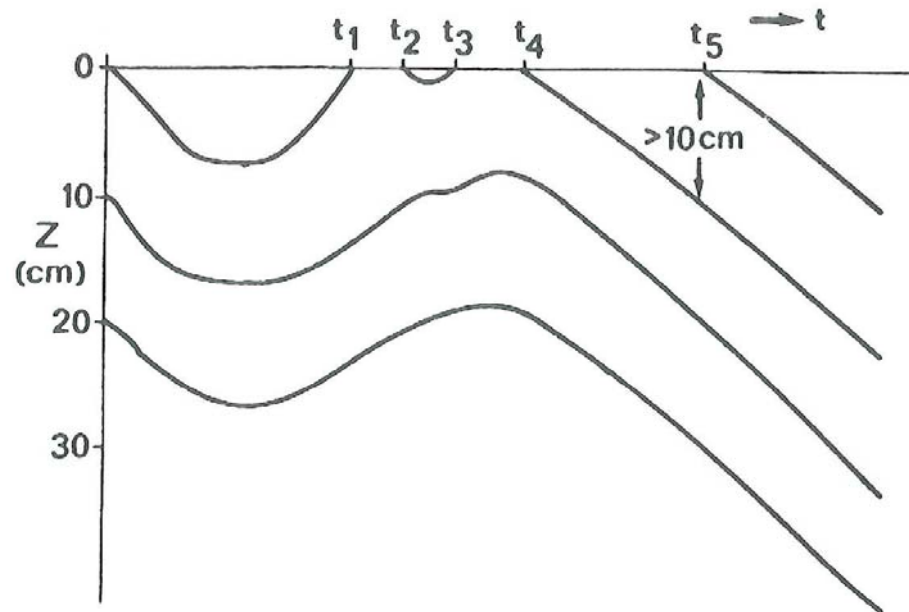


$$\int_0^{z_{t1}} \theta(z, t_1) dz = \int_0^{z_{t0}} \theta(z, t_0) dz + \int_{t_0}^{t_1} (N - E_a - E_i) dt - \int_{t_0}^{t_1} \int_0^{z_t} S(z, t) dz dt$$

time-depth-curves of a waterparcel



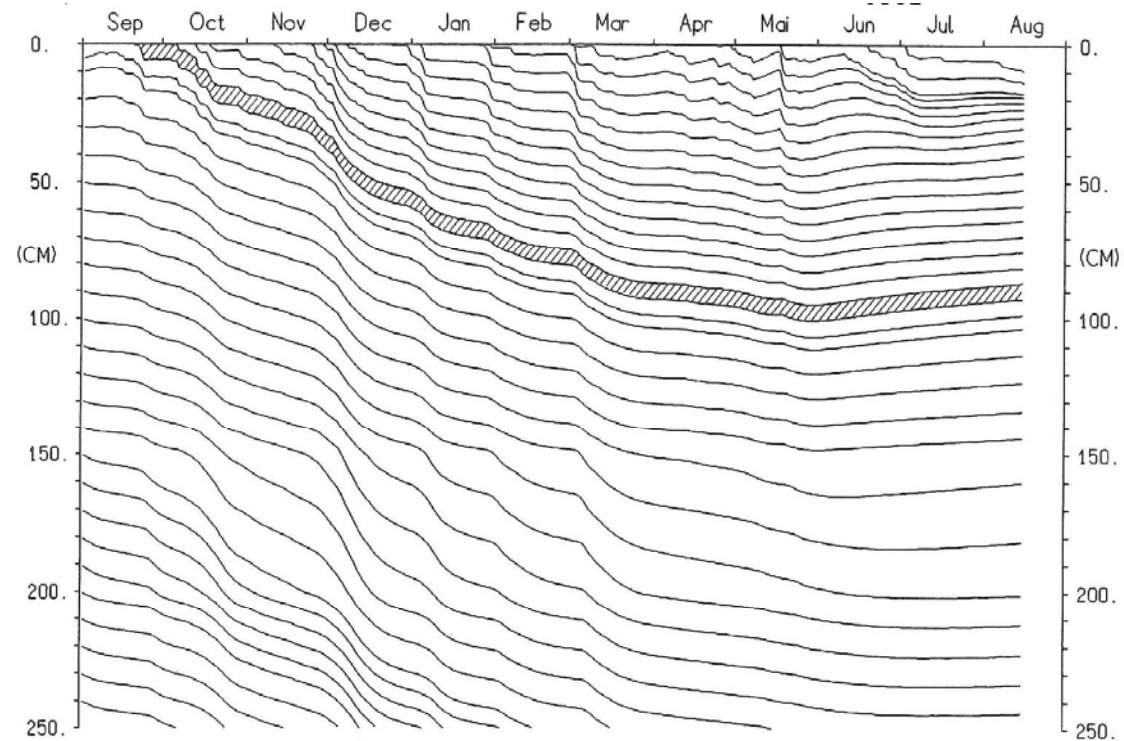
creation of new waterparcels:



- criteria:
- certain dates
 - amount of water in waterparcel
 - thickness of the waterparcel

Time-depth-curves of:

- Loess soil
- Summer wheat crop
- GW-table at 2.5 m
- Period: Sept - Aug



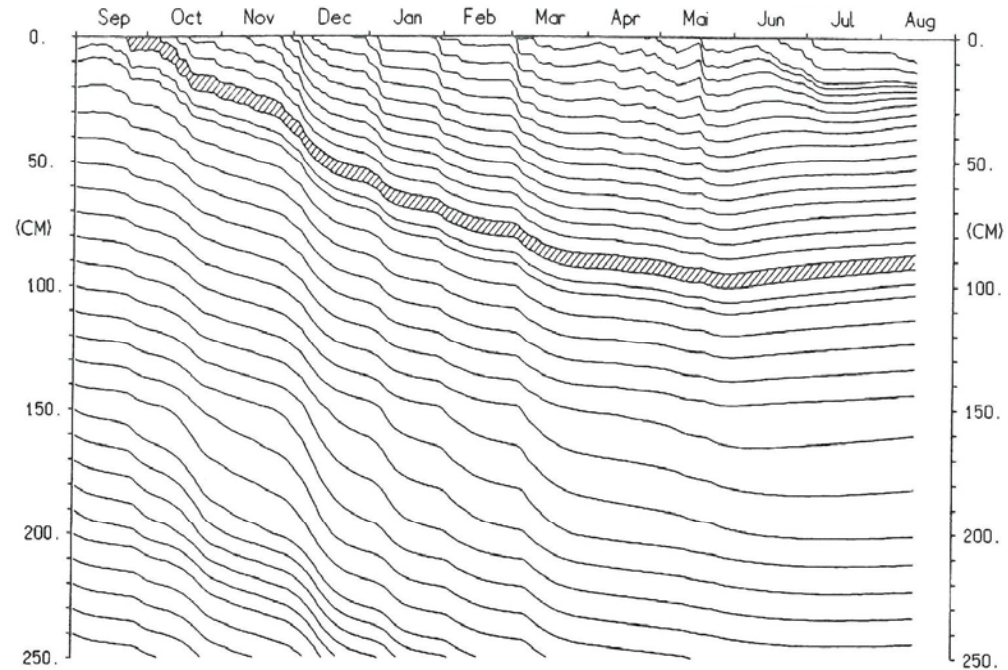
Water balance of :

B : WATER PARCELS PRESCRIBED AT THE BEGINNING OF THE SIMULATION PERIOD :

DEPTH-RANGE (T=0) CM	W (T=0) CM	W-UPT CM	W (T=T) CM	DEPTH-RANGE (T=T) CM
0.0 - 5.0	1.52	0.19	1.81	92.6 - 98.5
5.0 - 10.0	1.52	0.01	1.51	98.5 - 103.4
10.0 - 20.0	3.04	0.00	3.04	103.4 - 113.2
20.0 - 30.0	3.10	0.00	3.10	113.2 - 123.1
30.0 - 40.0	3.29	0.00	3.29	123.1 - 133.5
40.0 - 50.0	3.31	0.00	3.31	133.5 - 143.9
50.0 - 60.0	3.42	0.00	3.42	143.9 - 160.1
60.0 - 70.0	3.42	0.00	3.42	160.1 - 181.6

Time-depth-curves of:

- Loess soil
- Summer wheat crop
- GW-table at 2.5 m
- Period: Sept - Aug



Water balance of :

A : WATER PARCELS CREATED DURING THE SIMULATION PERIOD

	DATUM		* RAIN * CM	EI CM	EA CM	W-UPT CM	W(T=T) CM	DEPTH-RRAGE CM		
3. 7	205.1	- 15. 8	249.0	4.11	1.10	0.16	1.12	1.74	0.0	- 9.5
18. 6	190.8	- 3. 7	205.1	3.44	0.85	0.19	1.87	0.53	9.5	- 13.7
18. 5	159.5	- 18. 6	190.8	5.44	0.67	1.51	2.79	0.47	13.7	- 17.7
29. 4	140.3	- 18. 5	159.5	4.69	0.07	2.44	2.02	0.16	17.7	- 19.0
3. 3	83.9	- 29. 4	140.3	6.08	0.00	4.07	1.77	0.24	19.0	- 21.0
1. 3	82.0	- 3. 3	83.9	2.11	0.00	0.11	1.72	0.28	21.0	- 23.5
30. 1	51.7	- 1. 3	82.0	2.43	0.00	0.43	1.68	0.32	23.5	- 26.3
26. 1	47.2	- 30. 1	51.7	2.10	0.00	0.10	1.54	0.46	26.3	- 30.4
4. 1	25.4	- 26. 1	47.2	2.03	0.00	0.03	1.48	0.52	30.4	- 34.5
24.12	14.5	- 4. 1	25.4	2.11	0.00	0.11	1.28	0.72	34.5	- 40.2
7.12	-2.0	- 24.12	14.5	2.08	0.00	0.08	1.08	0.92	40.2	- 46.4
3.12	-6.1	- 7.12	-2.0	2.06	0.00	0.06	0.91	1.09	46.4	- 52.6
29.11	-10.3	- 3.12	-6.1	2.04	0.00	0.04	0.81	1.19	52.6	- 58.3
25.11	-15.0	- 29.11	-10.3	2.07	0.00	0.07	0.75	1.25	58.3	- 63.8
11.11	-28.9	- 25.11	-15.0	2.31	0.00	0.31	0.62	1.38	63.8	- 69.6
18.10	-52.5	- 11.11	-28.9	2.62	0.00	0.62	0.41	1.59	69.6	- 75.6
13.10	-57.5	- 18.10	-52.5	2.17	0.00	0.17	0.37	1.63	75.6	- 81.5
6.10	-64.7	- 13.10	-57.5	2.51	0.00	0.51	0.33	1.67	81.5	- 86.9
20. 9	-80.1	- 6.10	-64.7	3.09	0.00	1.09	0.27	1.73	86.9	- 92.6
0. 0	0.0	- 20. 9	-80.1	2.17	0.00	1.69	0.19	1.81	92.6	- 98.5

Notation for functions of several variables

Assume u is a function of the independent variables x and t . Subdivide the x - t plane into sets of equal rectangles of sides $\delta x = h$, $\delta t = k$, as shown in Fig. 1.4, and let the co-ordinates (x, t) of the representative mesh point P be

$$x = ih; \quad t = jk,$$

where i and j are integers.

Denote the value of u at P by

$$u_P = u(ih, jk) = u_{i,j}.$$

Then by equation (1.4),

$$\left(\frac{\partial^2 u}{\partial x^2}\right)_P = \left(\frac{\partial^2 u}{\partial x^2}\right)_{i,j} \approx \frac{u\{(i+1)h, jk\} - 2u\{ih, jk\} + u\{(i-1)h, jk\}}{h^2}.$$

i.e.

$$\left(\frac{\partial^2 u}{\partial x^2}\right)_{i,j} \approx \frac{u_{i+1,j} - 2u_{i,j} + u_{i-1,j}}{h^2}, \quad (1.8)$$

with a leading error of order h^2 . Similarly,

$$\left(\frac{\partial^2 u}{\partial t^2}\right)_{i,j} \approx \frac{u_{i,j+1} - 2u_{i,j} + u_{i,j-1}}{k^2}, \quad (1.9)$$

with a leading error of order k^2 .

With this notation the forward-difference approximation for $\partial u / \partial t$ at P is

$$\frac{\partial u}{\partial t} \approx \frac{u_{i,j+1} - u_{i,j}}{k}, \quad (1.10)$$

with a leading error of $O(k)$.

time

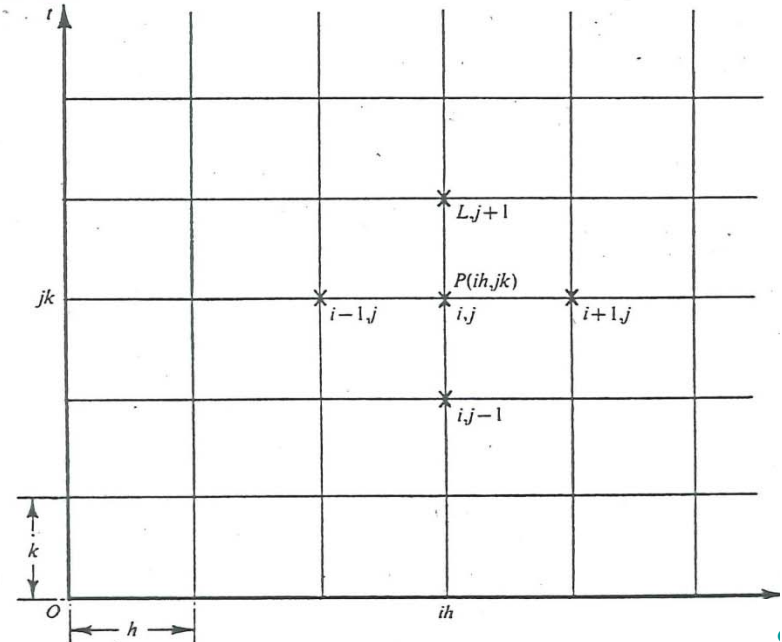


Fig. 1.4

$\delta x = h$

Finite-difference approximations to derivatives

When a function u and its derivatives are single-valued, finite and continuous functions of x , then by Taylor's theorem,

$$u(x+h) = u(x) + hu'(x) + \frac{1}{2}h^2u''(x) + \frac{1}{6}h^3u'''(x) + \dots \quad (1.1)$$

and

$$u(x-h) = u(x) - hu'(x) + \frac{1}{2}h^2u''(x) - \frac{1}{6}h^3u'''(x) + \dots \quad (1.2)$$

Addition of these expansions gives

$$u(x+h) + u(x-h) = 2u(x) + h^2u''(x) + O(h^4), \quad (1.3)$$

where $O(h^4)$ denotes terms containing fourth and higher powers of h . Assuming these are negligible in comparison with lower powers of h it follows that,

$$u''(x) = \left(\frac{d^2u}{dx^2}\right)_{x=x} \approx \frac{1}{h^2} \{u(x+h) - 2u(x) + u(x-h)\}, \quad (1.4)$$

with a leading error on the right-hand side of order h^2 .

Subtraction of equation (1.2) from equation (1.1) and neglect of terms of order h^3 leads to

$$u'(x) = \left(\frac{du}{dx}\right)_{x=x} \approx \frac{1}{2h} \{u(x+h) - u(x-h)\}, \quad (1.5)$$

with an error of order h^2 .

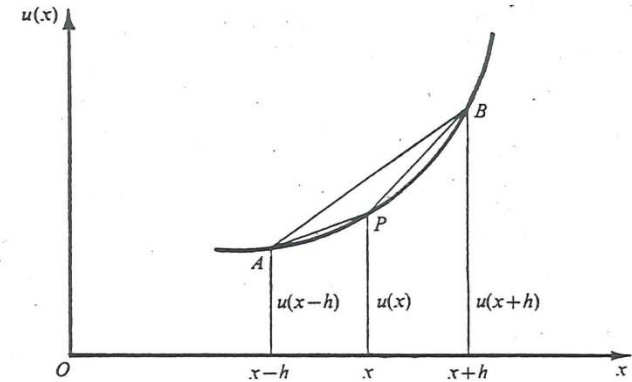


Fig. 1.3

Equation (1.5) clearly approximates the slope of the tangent by the slope of the chord AB , and is called a *central-difference* approximation. We can also approximate the slope of the tangent P by either the slope of the chord PB , giving the *forward-difference* formula,

$$u'(x) \approx \frac{1}{h} \{u(x+h) - u(x)\},$$

or the slope of the chord AP giving the *backward-difference* formula

$$u'(x) \approx \frac{1}{h} \{u(x) - u(x-h)\}.$$

Both (1.6) and (1.7) can be written down immediately from equations (1.1) and (1.2) respectively, assuming second and higher powers of h are negligible. This shows that the leading error in these forward and backward-difference formulae are both $O(h^2)$.